ROM SAF CDOP-2

Algorithm Theoretical Baseline Document: Level 2A dry temperature profiles

Version 1.1

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European Centre for Medium-Range Weather Forecasts (ECMWF)
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Met Office (METO)

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1.0	16/03/2016	SSY	Version for PCR-RE1 review. First version of ATBD for dry temperature
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ROM SAF

The Radio Occultation Meteorology Satellite Application Facility (ROM SAF) is a decentralised processing center under EUMETSAT which is responsible for operational processing of GRAS radio occultation (RO) data from the Metop satellites and radio occultation data from other missions. The ROM SAF delivers bending angle, refractivity, temperature, pressure, humidity, and other geophysical variables in near-real time for NWP users, as well as reprocessed data (Climate Data Records) and offline data for users requiring a higher degree of homogeneity of the RO data sets. The reprocessed and offline data are further processed into globally gridded monthly-mean data for use in climate monitoring and climate science applications.

The ROM SAF also maintains the Radio Occultation Processing Package (ROPP) which contains software modules that aids users wishing to process, quality-control and assimilate radio occultation data from any radio occultation mission into NWP and other models.

The ROM SAF Leading Entity is the Danish Meteorological Institute (DMI), with Cooperating Entities: i) European Centre for Medium-Range Weather Forecasts (ECMWF) in Reading, United Kingdom, ii) Institut D'Estudis Espacials de Catalunya (IEEC) in Barcelona, Spain, and iii) Met Office in Exeter, United Kingdom. To get access to our products or to read more about the ROM SAF please go to: http://www.romsaf.org

Intellectual Property Rights

List of Contents

1. Introduction

1.1 Purpose

This ATBD document describes the algorithms used to derive the dry temperature products produced by the Radio Occultation Meteorology (ROM) Satellite Application Facility (SAF). The complete list of products covered by this ATBD is provided in Table 1.1. The current operational status of all these and all other ROM SAF data products is available at the website: http://www.romsaf.org

The product requirements baseline is the PRD version 2.3 [AD.3]. The ATBD software package is based on the ROPP [RD.1].

Table 1.1 List of products covered by this ATBD

Product ID	Product name	Product acronym	Product type	Operational satellite input	Dissemination means	Dissemination format
GRM-29-L2-D-R1	Reprocessed Dry Temperature Profile	RDPMET	Data Record	Metop Level 1A data from EUM Secretariat	Web	BUFR/netCDF
GRM-30-L2-D-R1	Reprocessed Dry Temperature Profile	RDPCO1	Data Record	COSMIC Level 1A data from CDAAC	Web	BUFR/netCDF
GRM-32-L2-D-R1	Reprocessed Dry Temperature Profile	RDPCHA	Data Record	CHAMP Level 1A data from CDAAC	Web	BUFR/netCDF
GRM-33-L2-D-R1	Reprocessed Dry Temperature Profile	RDPGHA	Data Record	GRACE Level 1A data from CDAAC	Web	BUFR/netCDF

1.2 Applicable and reference documents

1.2.1 Applicable Documents

The following list contains documents with a direct bearing on the contents of this document:

[AD.1] CDOP-2 Proposal: Proposal for the Second Continuous Development and Operations Phase (CDOP-2); Ref: SAF/GRAS/DMI/MGT/CDOP2/001 Version 1.1 of 21 March 2011, approved by the EUMETSAT Council in Ref. EUM/C/72/11/DOC/10 at its 72nd meeting on 28-29 June 2011

[AD.2] CDOP-2 Cooperation Agreement: Agreement between EUMETSAT and DMI on the Second Continuous Development and Operations Phase (CDOP-2) of the Radio Occultation Meteorology Satellite Applications Facility (ROM SAF), approved by the EUMETSAT Council; Ref: EUM/C/72/11/DOC/15 at its 72nd meeting on 28-29 June 2011 and signed on 29 June 2011 in Copenhagen

[AD.3] ROM SAF CDOP-2 Product Requirements Document, Ref. SAF/ROM/DMI/MGT/PRD/001

1.2.2 Reference Documents

The following documents provide supplementary or background information, and could be helpful in conjunction with this document:

[RD.1] The Radio Occultation Processing Package (ROPP) User Guide, Part III: Pre-processor module, Ref. SAF/ROM/METO/UG/ROPP/004

[RD.2] Gorbunov ME (2009) Upgrading of OCC code for operational processing of GRAS raw sampling data. ROM SAF CDOP Visiting Scientist Report 6, Ref: SAF/GRAS/DMI/MGT/CVS06/003

[RD.3] Smith, E. K., and S. Weintraub (1953), The constants in the equation for the atmospheric refractive index at radio frequencies, Proc. IRE, 41(8), 1035–1037.

[RD.4] Aparicio, J. M., and S. Laroche (2011), An evaluation of the expression of the atmospheric refractivity for GPS signals, J. Geophys. Res., 116, D11104, doi:10.1029/2010JD015214

[RD.5] Gorbunov ME, Lauritsen KB, Rhodin A, Tomassini M, Kornblueh L (2006) Radio holographic filtering, error estimation, and quality control of radio occultation data. J. Geophys. Res. 111:D10105, doi:10.1029/2005JD006427

[RD.6] Algorithm Theoretical Baseline Document: Level 2A refractivity profiles, Ref. SAF/ROM/DMI/ALG/REF/001

[RD.7] Gorbunov ME, Shmakov AV, Leroy SS, Lauritsen KB (2011) COSMIC radio occultation processing: Cross-center comparison and validation. J Atmos Ocean Technol 28:737–751.

[RD.8] Algorithm Theoretical Baseline Document: Level 1B bending angles, Ref. SAF/ROM/DMI/ALG/BA/001

[RD.9] The Radio Occultation Processing Package (ROPP) User Guide, part I: Input/Output module, Ref. SAF/ROM/METO/UG/ROPP/002

[RD.10] Rüeger JM (2002), Refractive index formulae for electronic distance measurement with radio and millimeter waves. Tech. Rep. S-68, School of Surv. and Spatial Inf. Syst., Univ of N.S.W., Sydney, Australia

[RD.11] Healy SB (2011), Refractivity coefficients used in the assimilation of GPS radio occultation measurements. J Geophys Res 116, D01106, doi:10.1029/2010JD014013

[RD.12] Gorbunov ME (2002) Canonical transform method for processing radio occultation data in the lower troposphere. Radio Sci. 37:1076, doi:10.1029/2000RS002592

[RD.13] Lauritsen KB, Syndergaard S, Gleisner H, Gorbunov ME, Rubek F, Sørensen MB, Wilhelmsen H (2011) Processing and validation of refractivity from GRAS radio occultation data. Atmos. Meas. Tech. 4:2065-2071, doi:10.5194/amt-4-2065-2011

[RD.14] Syndergaard S (2012) Assessment of the Structural Uncertainty of GRAS Products from Level 1B (bending angles) up to Level 2 (temperatures), Final Report, Danish Meteorological Institute, EUMETSAT Contract No.

EUM/CO/10/4600000745/AvE.

[RD.15] Scherllin-Pirscher B, Syndergaard S, Foelsche U, and Lauritsen KB (2015) Generation of a bending angle radio occultation climatology (BAROCLIM) and its use in radio occultation retrievals. Atmos. Meas. Tech. 8:109-124, doi:10.5194/amt-8-109-2015.

[RD.16] Zeng Z, Sokolovskiy S (2010) Effect of sporadic E clouds on GPS radio occultation signals. Geophys Res Lett 37:L18817, doi:10.1029/2010GL044561

[RD.17] Sokolovskiy SV (2001) Tracking tropospheric radio occultation signals from low Earth orbit. Radio Sci 36:483-498

[RD.18] Sokolovskiy SV, Schreiner W, Rocken C, and Hunt D (2009) Optimal noise filtering for the ionospheric correction of GPS radio occultation signals. J. Atmos. & Oceanic Tech. 26:1398-1403, doi:10.1175/2009JTECHA1192.1

[RD.19] Algorithm Theoretical Baseline Document: Level 3 gridded data, Ref. SAF/ROM/DMI/ALG/GRD/001

1.3 Acronyms and abbreviations

ATBD	Algorithm Theoretical Baseline Document
BA	Bending Angle
BAROCLIM	Bending Angle Radio Occultation Climatology
CDAAC	COSMIC Data Analysis and Archive Center
CDOP-2	Second Continuous Development and Operations Phase
COSMIC	Constellation Observing System for Meteorology, Ionosphere, and Climate
CL	Closed Loop
DMI	Danish Meteorological Institute
ECF	Earth Centered Fixed
ECI	Earth Centered Inertial
ECMWF	European Center for Medium-range Weather Forecast
EDC	EUMETSAT Data Centre (former UMARF)
EGM96	Earth Gravitational Model
EPS	EUMETSAT Polar satellite System
EUMETSAT	EUropean organisation for the exploitation of METeorological SATellites
GMST	Greenwich Mean Sidereal Time
GNSS	Global Navigation Satellite System
GO	Geometric Optics
GPS	Global Positioning System (US)
GRAS	GNSS Receiver for Atmospheric Sounding (Metop instrument)
IEEC	Institut d'Estudis Espacials de Catalunya (Spain)
LEO	Low Earth Orbit
Metop	Meteorological Operational Polar satellite (EPS/EUMETSAT)
MSIS	Mass Spectrometer and Incoherent Scatter
MSL	Mean Sea Level
NCO	Numerically Controlled Oscillator

NIMA	National Imagery and Mapping Agency
NRT	Near-Real Time
NWP	Numerical Weather Prediction
OL	Open Loop
OLC	Optimal Linear Combination
RMS	Root Mean Square
RO	Radio Occultation
ROM SAF	Radio Occultation Meteorology SAF (EUMETSAT), former GRAS SAF
ROPP	Radio Occultation Processing Package
RS	Raw Sampling
SAF	Satellite Application Facility (EUMETSAT)
SLTP	Straight-Line Tangent Point
SNR	Signal-to-noise ratio
UKMO	The UK Meteorological Office (aka: Met Office)
WGS84	World Geodetic System
WMO	World Meteorological Organisation
WO	Wave Optics

1.4 Definitions

RO data products from the GRAS instrument onboard Metop and RO data from other missions are grouped in *data levels* (level 0, 1, 2, or 3) and *product types* (NRT, Reprocessed, or Offline). The data levels and product types are defined below. The lists of variables should not be considered as the complete contents of a given data level, and not all data may be contained in a given data level.

Data levels:

Level 0: Raw sounding, tracking and ancillary data, and other GNSS data before clock correction and reconstruction;

Level 1A: Reconstructed full resolution excess phases, total phases, pseudo ranges, SNR's, orbit information, I, Q values, NCO (carrier) phases, navigation bits, and quality information;

Level 1B: Bending angles and impact parameters, tangent point location, and quality information;

Level 2: Refractivity, geopotential height, "dry" temperature profiles (level 2A), pressure, temperature, specific humidity profiles (level 2B), surface pressure, tropopause height, planetary boundary layer height (level 2C), ECMWF model level coefficients (level 2D); quality information;

Level 3: Gridded or resampled data, that are processed from level 1 or 2 data, and that are provided as, e.g., daily, monthly, or seasonal means on a spatiotemporal grid, including metadata, uncertainties and quality information.

Product types:

NRT product: Data product delivered less than 80 min (95%; EPS-SG Global), 40 min (95%; EPS-SG Regional) and 3 hours (EPS) after measurement;

Reprocessed product: Climate Data Record (CDR) covering an extended time period of several years, generated using a fixed set of processing software in order to provide a homogeneous data record appropriate for climate usage;

Offline product: Data product which typically extends a CDR from a reprocessing, thus providing an “interim” CDR; delivered from less than 5 days to up to 6 months after measurement depending on the requirements.

2. Algorithm overview

RO data may potentially have benchmarking quality for climate analyses because of the all-weather capability of the technique and because there is no need for calibration (as opposed to many other remote sensing instruments). However, RO processing is generally complex, not the least because different RO missions have different problems (such as low SNR, poor L2 tracking, data gaps, spikes, etc). Thus, besides the processing steps that can be easily described by equations, it is necessary to also have algorithms that can cope with a number of problematic issues. The algorithms in the Radio Occultation Processing Package (ROPP) have been developed over many years to do just that.

ROPP contains a module designed to compute ionospheric corrected bending angle, refractivity, and dry temperature profiles either from excess phase or L1 and L2 channel bending angle data measured during a radio occultation. A flow chart illustrating the ROPP pre-processor module is given in Figure 2.1. The main aspects of the algorithm for the level 2A dry temperature are described in the ROPP pre-processor user guide [RD.1].

The algorithm description in this ATBD complements the ROPP user guide by focusing on details not described in the user guide or elsewhere. References to equations and sections in the user guide are provided when appropriate. Many of the algorithms in the ROPP pre-processor module are also described in [RD.2]. References to original work on which algorithms are based are provided in the relevant sections.

In the descriptions in the following, the specific choices of parameters that affect the outcome of the processing is mentioned, such as filter widths, intermediate and output vertical grids, limits and parameters determining specifics in the processing at various stages, interpolation methods, etc. The values mentioned in this ATBD are the values that are either hard-coded in the software or set in a configuration file in the ROM SAF processing. Although these choices have influence on the results, and contribute to the structural uncertainty of the products, they are not considered to have any negative impact on the products and they do not compromise the benchmarking quality of the data.

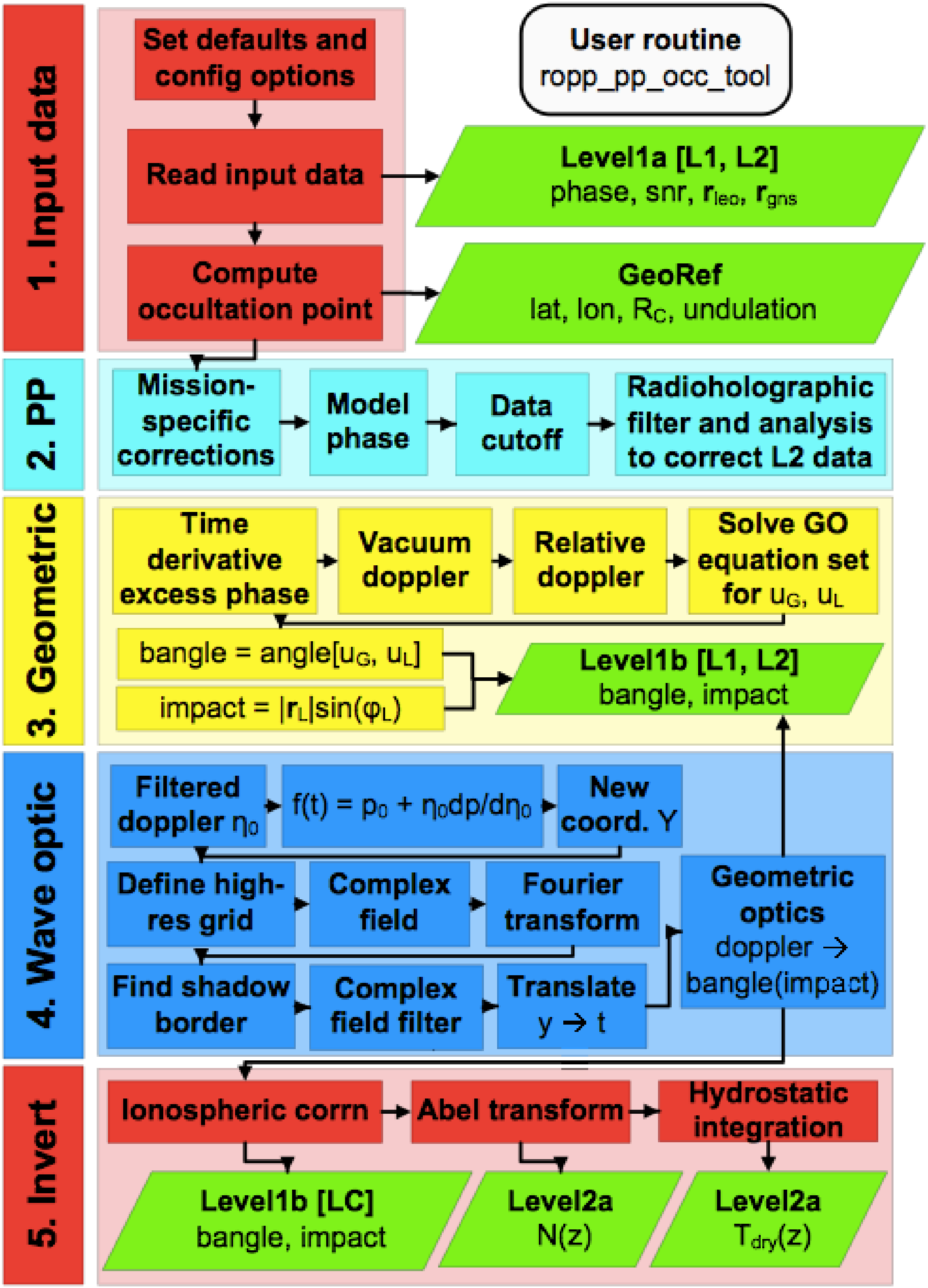

Figure 2.1: Flow chart illustrating calling tree of the ROPP pre-processor occ tool to compute ionospheric corrected bending angle, refractivity, and dry temperature profiles from input L1 and L2 channel amplitude and phase measurements [RD.1].

3. Algorithm description

3.1 Physics of the problem

3.1.1 Fundamental observables

The fundamental observables measured by an RO instrument are the phase, L_i, and amplitude, A_i, of the Doppler-shifted incoming signal. Index i denotes one of the two GNSS frequencies L1 (1575.42 MHz) and L2 (1227.60 MHz). Each occultation measurement is a time-series of measured phases and amplitudes as well as precise position information for the transmitter (GNSS) satellite and the receiver (LEO) satellite.

3.1.2 Doppler-shift and derived quantities

The received signal will be Doppler-shifted due to the motion of the transmitter and receiver satellites. With known satellite positions and velocities this Doppler-shift may be calculated to high precision for the vacuum case. When the ray bends in the atmosphere the angles between the ray path and the directions of motion change both for the transmitting and the receiving satellite. This leads to a change in the observed Doppler-shift. From observed signal phases the observed Doppler-shift may be found and from this the bending of the ray path through the atmosphere may be derived. This leads to a profile of bending angles as a function of impact parameter. Via the Abel transform the profile of bending angles as a function of impact parameter is converted to a profile of atmospheric refractive index as a function of altitude. For convenience results are given in terms of the refractivity N instead of the refractive index n, with refractivity defined by $N = 10^6 \cdot (n\text{-}1)$.

3.1.3 Relation to dry temperature

The refractivity is related to pressure, temperature, and humidity [RD.3; RD.4]. Assuming dry air and hydrostatic equilibrium, a temperature profile can be derived using the equation of state. Thus, we define $N = \kappa_1 \cdot P_{\mathrm{dry}}/T_{\mathrm{dry}}$, where P_{dry} is called the dry pressure, T_{dry} is called the dry temperature, and $\kappa_1 = 77.6$ N-unit·K·hPa^{-1}.

3.2 Mathematical description of the algorithm

Given a profile of refractivity as a function of altitude above the geoid, the following subsections describe the steps taken to obtain the dry temperature as a function of altitude above the geoid.

3.2.1 Hydrostatic integration

The dry temperature (and corresponding dry pressure) is obtained by ignoring the water vapor contribution to refractivity [RD.5]. Using the equation of state for an ideal gas and assuming hydrostatic equilibrium, the equation to be solved can be written (Eq. 4.20 in [RD.1]):

$$\frac{d \ln P_{\mathrm{dry}}}{dz} = f(z, \ln P_{\mathrm{dry}}(z)) = -\frac{g(z)N(z)}{R\kappa_1 \exp(\ln P_{\mathrm{dry}}(z))},$$

where $g(z)$ is the gravitational acceleration as a function of altitude, z, and R is the dry air gas constant. The dry pressure at each level is obtained using a fourth order Runge-Kutta method starting from the top at 150 km. The dry temperature readily follows as $T_{\text{dry}} = \kappa_1 \cdot P_{\text{dry}}/N$. The integration step of the Runge-Kutta method is ~15 m. During integration the refractivity is interpolated (using spline interpolation of log(N)) to the high-resolution levels. After the integration, temperature and pressure are interpolated back to the original levels (~100 m spacing).

3.2.2 Upper boundary condition

For initialization of the hydrostatic integration, the pressure at 150 km is calculated from the refractivity gradient at this altitude, disregarding the temperature gradient (Eq. 4.21 in [RD.1]).

3.3 Error sources

As a general consideration, the radio occultation signal consists of an excess phase and an amplitude. High-quality data is data with high SNR both in terms of amplitude and in terms of excess phase.

Amplitude noise is dominated by instrument noise under quiet ionospheric conditions. However, in the presence of ionospheric disturbances, or tilted sporadic E-layers, it can be severely affected by scintillations [RD.16]. Under quiet conditions the SNR is generally high except in the middle to lower troposphere, where the denser atmosphere leads to loss of signal intensity. This is particularly true when the humidity is high, which typically occurs in the tropics [RD.17]. This results in degraded bending angle data quality in the lower to middle troposphere, particularly in the tropics.

Besides instrument noise, the measured excess phase is affected by residual ionospheric noise [RD.18]. The ionospheric contribution to the signal is not fully removed by the linear combination of the L1 and L2 signals due to short timescale ionospheric variation and other higher order effects. Since the measured excess phase signal is a function of atmospheric density it falls off approximately exponentially with impact height and so the noise comes to dominate the signal at high altitudes in the upper stratosphere and above. In the lower troposphere the tracking of the L2 signal becomes difficult and for that reason only the L1 signal is useful. This limits the accuracy of derived products.

The highest data quality is therefore found at intermediate altitudes of the higher troposphere to lower stratosphere, where the signal is strong both in terms of amplitude and in terms of excess phase.

Above a certain altitude in the upper stratosphere to lower mesosphere (depending on the noise level) the retrieved dry temperature is based on a merge between observations and a climatology (see also [RD.19]). The implementation of the global search and fitting procedure in the statistical optimization of bending angles to get to the refractivity (see [RD.6]) should ensure a very little bias influence from the climatology.

4. Practical considerations

4.1 Validation method

As a whole, the algorithms will be used to process a number of occultation observations, which are then compared to the corresponding profiles extracted from ECMWF analyses and forecasts (forward modelled to dry temperature as a function of altitude above the geoid). The dry temperature profiles based on input data from CDAAC will also be compared to the corresponding dry temperature profiles produced by CDAAC.

As the dry temperature is derived from the refractivity, which in turn is derived from the bending angle, the validation of all parts of the processing chain from bending angle to dry temperature is relevant to ensure the quality of the dry temperature. Many parts of the algorithms described here together with those described in [RD.6] and [RD.8], have been validated over many years, as similar versions of the algorithms have been used to produce results for scientific publications and reports (see [RD.2], [RD.5], [RD.7], [RD.12], [RD.13], and [RD.14]).

Certain parts have been modified over the past few years in the version of ROPP at DMI. These parts include an improvement to the search and fitting strategy to find a suitable background for the statistical optimization and the development and inclusion of the BAROCLIM model (Section 3.2.4 in [RD.6]). These modifications were validated by comparisons to data produced by the unmodified code, comparisons against ECMWF analyses, and comparisons to corresponding profiles from CDAAC. The algorithms have also been validated by comparing ‘raw’ and optimized bending angles. The latter approach is an efficient way to evaluate to which extent a potential bias from the climatology affects the retrievals. The generation and the validation of BAROCLIM for its use in radio occultation retrievals (though with a different search and fitting strategy than the one used here) can be found in [RD.15].

4.2 Quality control and diagnostics

The following quality control parameters are used to ensure the quality of the dry temperature products:

L2 quality score:

> Measures the quality of the L2 signal. This score is defined as the maximum of an L2 penalty function over the interval 15 km - 50 km (see [RD.8]).

SO quality score:

> This is based on the estimated profile of error variances of the bending angle solution (see section 2(e) of [RD.7]), and is estimated after the performance of statistical optimization. This variance estimate has two terms: One term is from the statistical optimization and ionospheric correction procedure and measures the quality of the match between the data and the fitted background profile. The second term is a radio-holographic error variance on the L1 signal (from the CT2 wave-optics processing; see [RD.8]), which is a measure of the noisiness of the L1 signal. The SO quality score is constructed as the maximum over the entire profile of the

square root of the solution bending angle variance estimate taken as a percentage of the solution bending angle.

Both quality scores are constructed such that a low value means high quality data.

Scaling factors:

The fitting of the background profile to the data at high altitudes results in two scaling factors. Usually the scaling factors are close to unity, but in cases of erroneous data, they can be off by large factors, and are therefore used as an additional quality check.

Quality scores and scaling factors are generated at different places in the code when the relevant parameters to generate them are readily available. The quality scores and scaling factors are output together with the data and are common to bending angle, refractivity, and dry temperature products. If any of the quality scores or scaling factors exceeds defined thresholds, all three products are considered to have poor quality and marked as non-nominal.

As a final sanity check, the dry temperature is compared to the corresponding profile extracted from the ECMWF analysis (forward modeled to dry temperature as a function of altitude above the geoid). If the dry temperature fails this sanity check (if the difference is outside defined thresholds), the level 2A data is marked as non-nominal.

4.3 Exception handling

N/A

4.4 Outputs

The output of the processing to dry temperature is a ROPP NetCDF file containing the following profile variables:

- Altitude above the geoid
- Geopotential height
- Dry temperature

The same NetCDF file contains the output from the bending angle [RD.8] and refractivity [RD.6] processing. A more complete and technical description of the output to the NetCDF file can be found in [RD.9].

5. Assumptions and limitations

5.1 Assumptions

5.1.1 Spherical symmetry

Radio occultation data are generally processed under the assumption of spherical symmetry. However, in principle this is only an apparent assumption because it depends on the interpretation of the retrieved profiles. If profiles are interpreted as representing the vertical structure in the atmosphere at a given fixed location, then the spherical symmetry assumption gives rise to a real error because the atmosphere is only approximately spherically symmetrical. If, on the other hand, retrieved profiles are interpreted as being weighted averages of the 3-dimensional (3D) atmosphere (primarily in the 2-dimensional (2D) occultation plane), the spherical symmetry assumption does not in principle give rise to any errors. This is why it could be an advantage to assimilate occultation data with 2D or 3D observation operators. Although the assimilation of dry temperature would be possible (with a suitable observation operator), it is not used in any known operational forecasting systems, and thus for all practical use, the spherical symmetry assumption gives rise to errors in the dry temperature.

5.1.2 The assumption of no water vapor

The dry temperature is not the actual (thermodynamic) temperature, but for all practical purposes it is the same as the actual temperature at altitudes where water vapour gives a negligible contribution to the refractivity (in practice above ~10 km depending on latitude; a rule of thumb is that the water vapour is negligible where the temperature is less than 240 K or at altitudes above the tropopause). However, even in the lower troposphere, where water vapour is normally abundant, the dry temperature can be considered a new useful variable on its own, just as we consider the better-known virtual temperature a variable on its own. Thus the assumption of no water vapour can be considered just a way to define dry temperature.

5.1.3 Equation of state for an ideal gas

In the retrieval of dry temperature from refractivity it is assumed that the atmosphere is an ideal gas, which makes refractivity proportional to density. This is a very good assumption for most practical applications. Careful studies on the general accuracy of the relation between refractivity, pressure, temperature and water vapour can be found in [RD.4], [RD.10], and [RD.11]. There is not a unique conclusion on the best or most accurate values for the coefficients, and the issue is complicated by the fact that the coefficients seem to vary with atmospheric conditions.

5.1.4 Hydrostatic equilibrium

The dry pressure, and thus dry temperature, is obtained using the assumption of hydrostatic equilibrium. This is a very good assumption for most practical applications.

5.2 Algorithm limitations

The following subsections discuss limitations in the algorithms described in the corresponding subsections with the same titles in Section 3.2.

5.2.1 Upper boundary condition

In very rare cases the retrieved refractivity at very high altitudes is slightly negative (presumably related to the dynamical error estimation and ionospheric scintillations; see [RD.6]), which prevents further processing to dry temperature.

6. Description of differences for NRT, Offline and Reprocessing

This chapter describes the parts of the algorithm which are different for NRT, Offline and Reprocessed products.

6.1 NRT

The algorithms used in NRT are the same as the algorithms used in offline and reprocessing.

6.2 Offline

The algorithms used in offline are the same as the algorithms used in NRT and reprocessing.

6.3 Reprocessing

The algorithms used in reprocessing are the same as the algorithms used in NRT and offline.

7. Appendices

7.1 Description of how to run the code

The code is run by the following command (for offline and reprocessing):

```
ropp_pp_occ_tool <input_file> --no-ranchk -o <output_file> -c
<config_file>
```

The input file is a ROPP NetCDF file containing high-resolution level 1a data. The output file is a ROPP NetCDF file containing high-resolution level 1b and 2a data. An example of a configuration file is given in Section 7.2.

7.2 Configuration files

An example of a ROPP PP configuration file is given below. The values of parameters are not necessarily the final ones that will be set in the offline and reprocessing of dry temperature.

```
# $Id: $

#****c* Configuration Files/cosmic_pp.cf *
#
# NAME
#    default_pp.cf - COSMIC data configuration file for pre-processor
#                    implementations in ROPP
#
# SYNOPSIS
#    <pp_program> ... -c cosmic_pp.cf ...
#
# DESCRIPTION
#    This file reflects the configuration for the PP
#    implementations within ROPP suitable for use with COSMIC data.
#
# NOTES
#
# AUTHOR
#   Met Office, Exeter, UK.
#   Any comments on this software should be given via the ROM SAF
#   Helpdesk at http://www.romsaf.org
#
# COPYRIGHT
#   (c) EUMETSAT. All rights reserved.
#   For further details please refer to the file COPYRIGHT
#   which you should have received as part of this distribution.
#
#****

#--------------------------------------------------------------------------------
# 0. Output options
#--------------------------------------------------------------------------------
output_lev1a = .false.      # Flag to output (modified) level 1a data

output_lev1b = .true.       # Flag to output level 1b data

output_lev2a = .true.       # Flag to output level 2a data

output_diag = .false.       # Flag to output additional diagnostics

#--------------------------------------------------------------------------------
```

```
# 1. Excess phase to bending angle processing
#-------------------------------------------------------------------------------

# 1.1 Occultation processing method
# ---------------------------------

# GO - use GEOMETRIC OPTICS processing to derive bending angle as a function of
#      impact parameter from excess phase as a function of time.
# WO - use WAVE OPTICS (CT2 algorithm) processing to derive bending angle as a
#      function of impact parameter from excess phase as a function of time.

occ_method = WO

# 1.2 Filtering method
# --------------------

# optest - use OPTIMAL ESTIMATION: solution of integral equation
# slpoly - use SLIDING POLYNOMIAL

filter_method = slpoly

# 1.3 Smoothing bending angle profile
# -----------------------------------


fw_go_smooth = 3000.0   # Filter width for smoothed GO bending angles (m)

fw_go_full  = 3000.0    # Filter width for full resolution GO bending angles (m)

fw_wo = 2000.0           # Filter width for wave optics bending angle above 7
km(m)

fw_low = -1000.0           # Filter width for wave optics bending angle below 7
km (m)

# 1.4 Maximum height for wave optics processing
# ---------------------------------------------

hmax_wo = 25000.0      # Maximum height for wave optics processing (m)

# 1.5 Data cut-off limits
# -----------------------

Acut     = 0.0        # Fractional cut-off limit for amplitude

Pcut     = -2000.0        # Cut-off limit for impact height

Bcut     = 0.1       # Cut-off limit for bending angle

Hcut     = -250000.0     # Cut-off limit for straight-line tangent altitude

# 1.6 CT2 options
# ---------------

CFF      = 3        # Complex field filter flag (CFF = 'Pa')

dsh      = 200.0    # Shadow border width (m)

# 1.7 Degraded L2 data flag
# -------------------------

opt_DL2  = .true.

# 1.8 Compute and output spectra flag
# -----------------------------------
```

```
opt_spectra = .false.

# 1.9 Paths to EGM96 geoid model coefficients and corrections file
# ----------------------------------------------------------------

egm96 = ../data/egm96.dat                # EGM96 coefficients file

corr_egm96 = ../data/corrcoef.dat        # Correction coefficients file

#-------------------------------------------------------------------------------
# 1. Ionospheric correction processing
#-------------------------------------------------------------------------------

# 1.1 Ionospheric correction method
# ---------------------------------

# GMSIS - use MSIS climatology bending angle (searching global MSIS profiles
#         for best fit profile to obs) in ionospheric correction,
#         statistical optimization and bending angle to refractivity inversion.
#
# MSIS - use MSIS climatology bending angle in ionospheric correction,
#        statistical optimization and bending angle to refractivity inversion.
#
# GBARO - use BAROCLIM bending angle (searching global BAROCLIM profiles
#         for best fit profile to obs) in ionospheric correction,
#         statistical optimization and bending angle to refractivity inversion.
#
# BARO - use BAROCLIM bending angle in ionospheric correction,
#        statistical optimization and bending angle to refractivity inversion.
#
# BG   - use climatology from a specified input file containing
#        background temperature, pressure and humidity
#        (e.g. from an NWP analysis). The input filename can be specified
#        using the '-bfile' command line argument or setting 'bfile' (see 1.5).
#
# NONE - linear combination of L1 and L2 bending angles in ionospheric
#       correction, no additional information above observed profile top
#        in the inverse Abel to compute refractivity.

method  = GMSIS                # Ionospheric correction method

# 1.2 Abel integral method
# ------------------------

# LIN  - assume linear variation of bending angle and ln(n) between
#        observation levels. This algorithm is used in ROM SAF NRT processing
#
# EXP  - assume exponential variation of bending angle and ln(n) between
#        observation levels. This algorithm is used in ropp_fm module.

abel = LIN

# 1.3 Statistical optimisation method
# -----------------------------------

# SO   - statistical optimisation.
# LCSO  - linear combination plus statistical optimisation.

so_method = so

# 1.4 Climatology model coefficients files
# --------------------------------

msisfile = MSIS_coeff.nc       # MSIS model coefficients file
```

```
mfile    = MSIS_coeff.nc       # Model coefficients file for stat.opt.

# 1.5 Background model temperature, humidity, pressure file
# ---------------------------------------------------------

bfile    = BG_file.nc       # Background meteorology profile file (method=BG)

#-------------------------------------------------------------------------------
# 2. Impact parameter grid
#-------------------------------------------------------------------------------

# The ionospheric correction interpolates L1 and L2 bending angle profiles onto a
# standard grid.

dpi = 100.0         # Step of standard impact parameter grid (m)

#-------------------------------------------------------------------------------
# 3. Smoothing bending angle profile
#-------------------------------------------------------------------------------

# A smoothed bending angle profile is derived compute the fit of observed bending
# angles to the model bending angle profile.

np_smooth = 3       # Polynomial degree for smoothing regression

fw_smooth = 1000.0      # Filter width for smoothing profile

#-------------------------------------------------------------------------------
# 4. Model bending angle profile fit to observations
#-------------------------------------------------------------------------------

# To avoid systematic deviations from the observed profile with climatology,
# the model profile is scaled to the observed profile by a fitting method.

sf_method = convoluted  # Search and fit method (convoluted or regular)

nparm_fit = 2           # Number of parameters for model fit regression

hmin_fit  = 20000.0     # Lower limit for model fit regression

hmax_fit  = 70000.0     # Upper limit for model fit regression

omega_fit = 0.3         # A priori standard deviation of regression factor

#-------------------------------------------------------------------------------
# 5. Ionospheric correction and statistical optimization
#-------------------------------------------------------------------------------

# The method described by Gorbunov (2002) is implemented to perform ionospheric
# correction with statistical optimization.

f_width = 2000.0    # Ionospheric correction filter width

delta_p = 20.0             # Step of homogeneous impact parameter grid

s_smooth = 2000.0   # External ionospheric smoothing scale

z_ion = 50000.0            # Lower height limit of ionospheric signal

z_str = 35000.0            # Lower height limit of stratospheric signal

z_ltr = 12000.0            # Lower height limit of tropospheric signal

n_smooth = 11       # Number of points for smoothing (must be odd)
```

```
model_err = 0.5         # A priori model error std.dev. (dyn.est. if negative)

#--------------------------------------------------------------------------------
# 6. Bending angle inversion to refractivity
#--------------------------------------------------------------------------------

# The Abel inversion is computed to retrieve refractivity from corrected
# bending angles. The corrected bending angle profile is extended
# using MSIS or BAROCLIM data above the observed profile top.

ztop_invert = 150000.0     # Height of atmosphere top for inversion

dzh_invert = 50.0   # Step of inversion grid above observation top

dzr_invert = 20000.0      # Interval for regression in inversion

#--------------------------------------------------------------------------------
# 7. Tangent point lat-lons
#--------------------------------------------------------------------------------

# Set tp_bending=.true. to update lat-lons accounting for bending

tp_bending = .false.
```